

Why is restored peatland NEE so high?

Insights from three methods for CO₂ flux estimates



Lisamarie Windham-Myers¹, Brian Bergamaschi² and Frank Anderson²

US Geological Survey, Menlo Park and Sacramento, CA

Freshwater marsh restoration on subsided Delta islands

(Peat fills the consistent 25-55cm accomodation space)



Estimates of NEE fluxes in San Francisco Bay-Delta lands

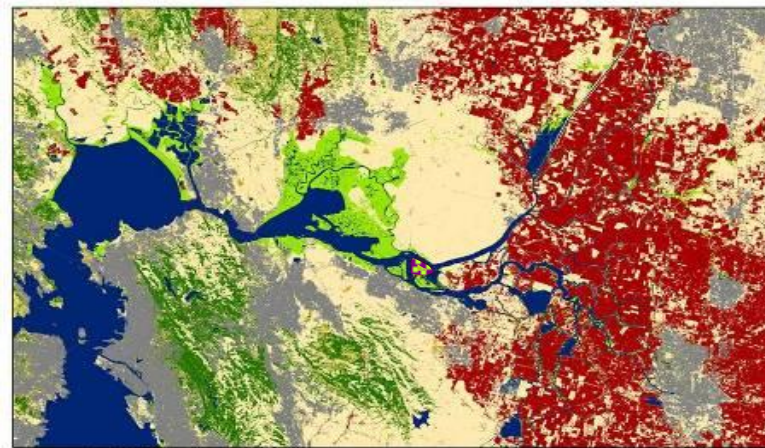
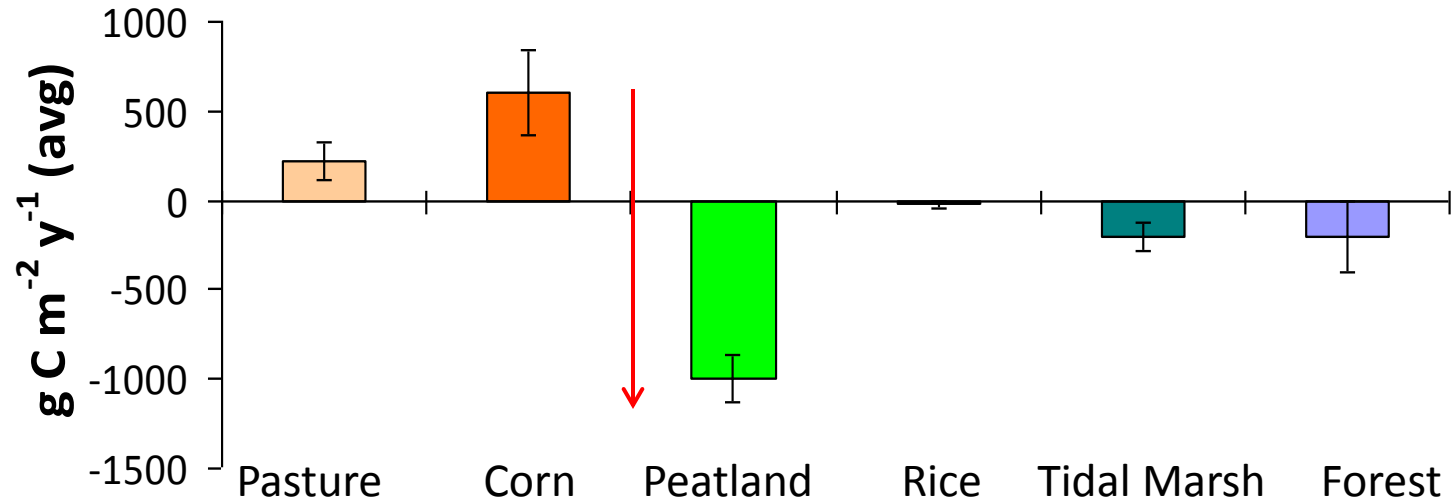


Image by Maggi Kelly, UC Berkeley

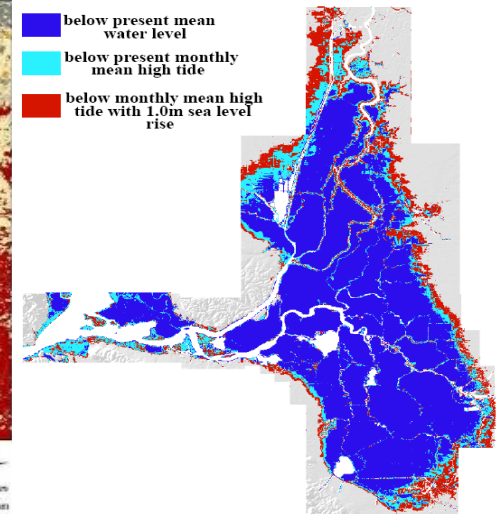
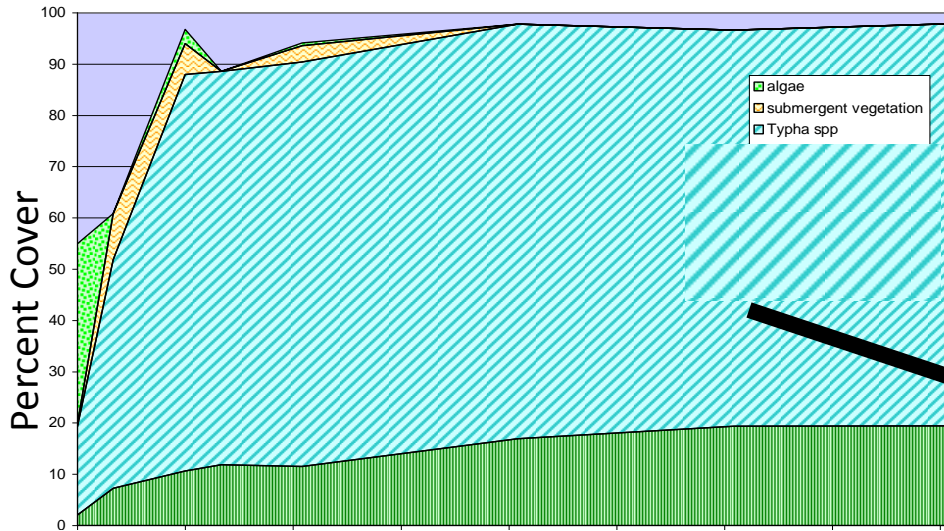


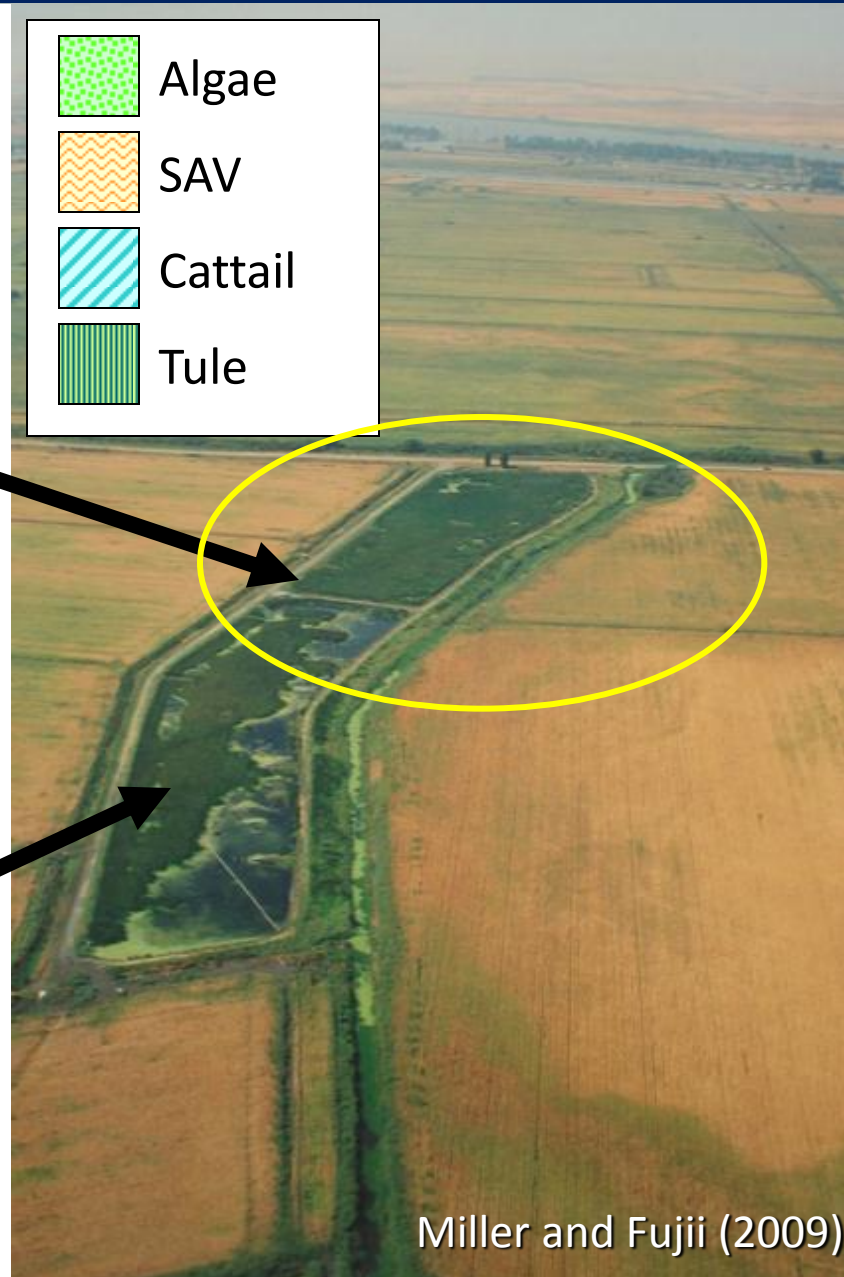
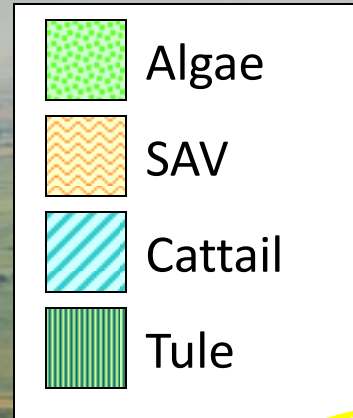
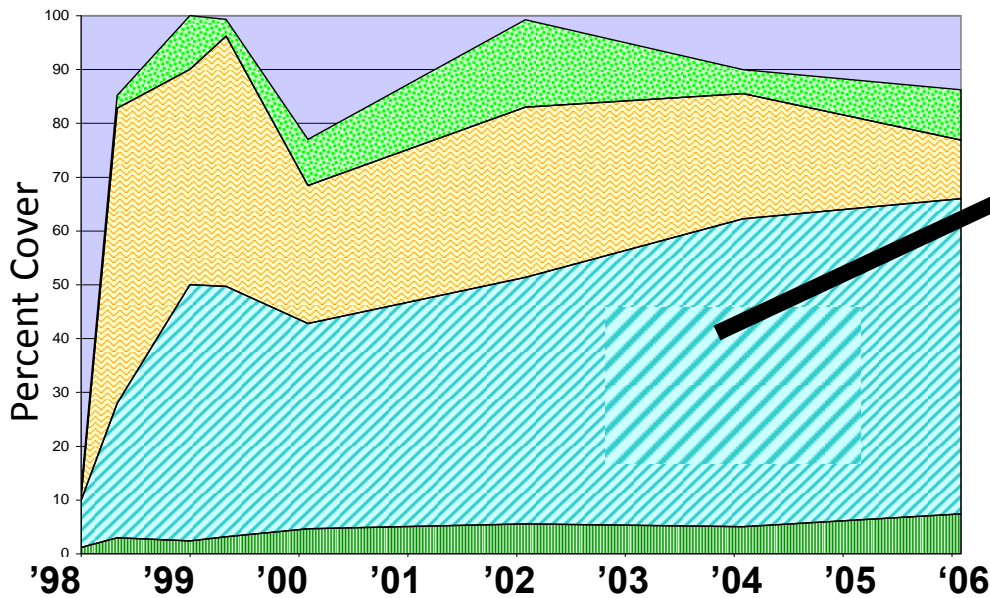
Image by Noah Knowles, USGS

West Wetland (25 cm depth) had best coverage

SHALLOW, West Wetland (25 cm)



DEEP, East Wetland (55 cm)



Three Methods (August 29-Sept 2, 2011)

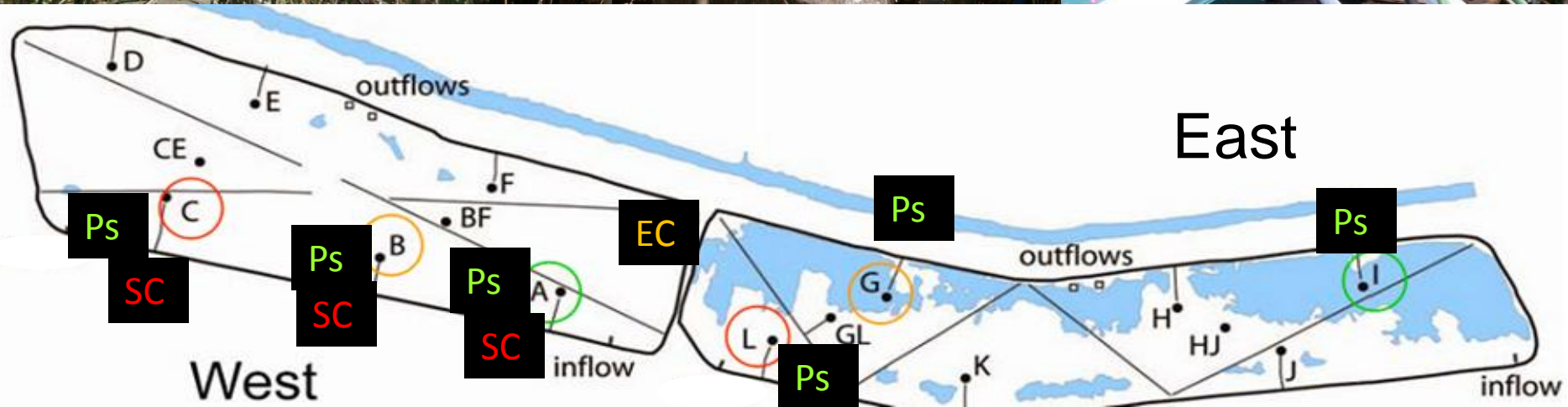
Eddy Covariance



Static Chamber



Leaf Photosynthesis



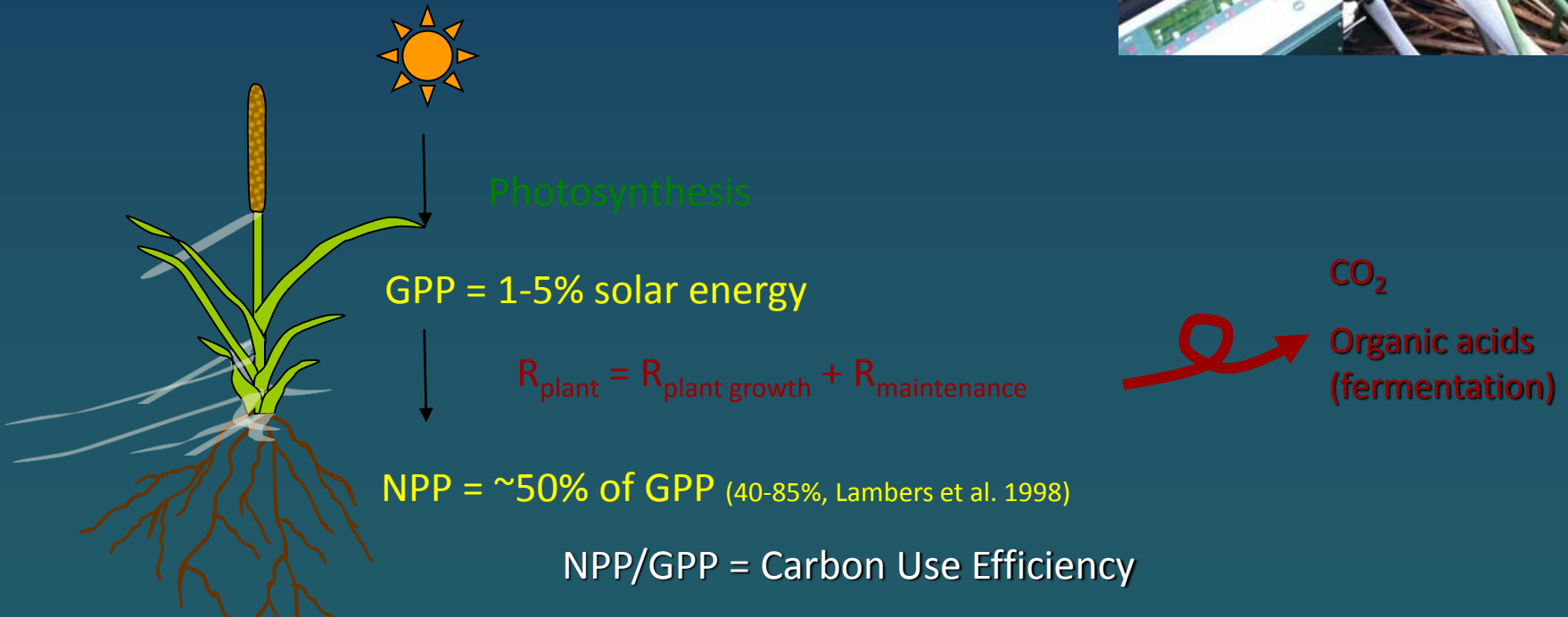
Leaf Photosynthesis

PHOTOSYNTHESIS

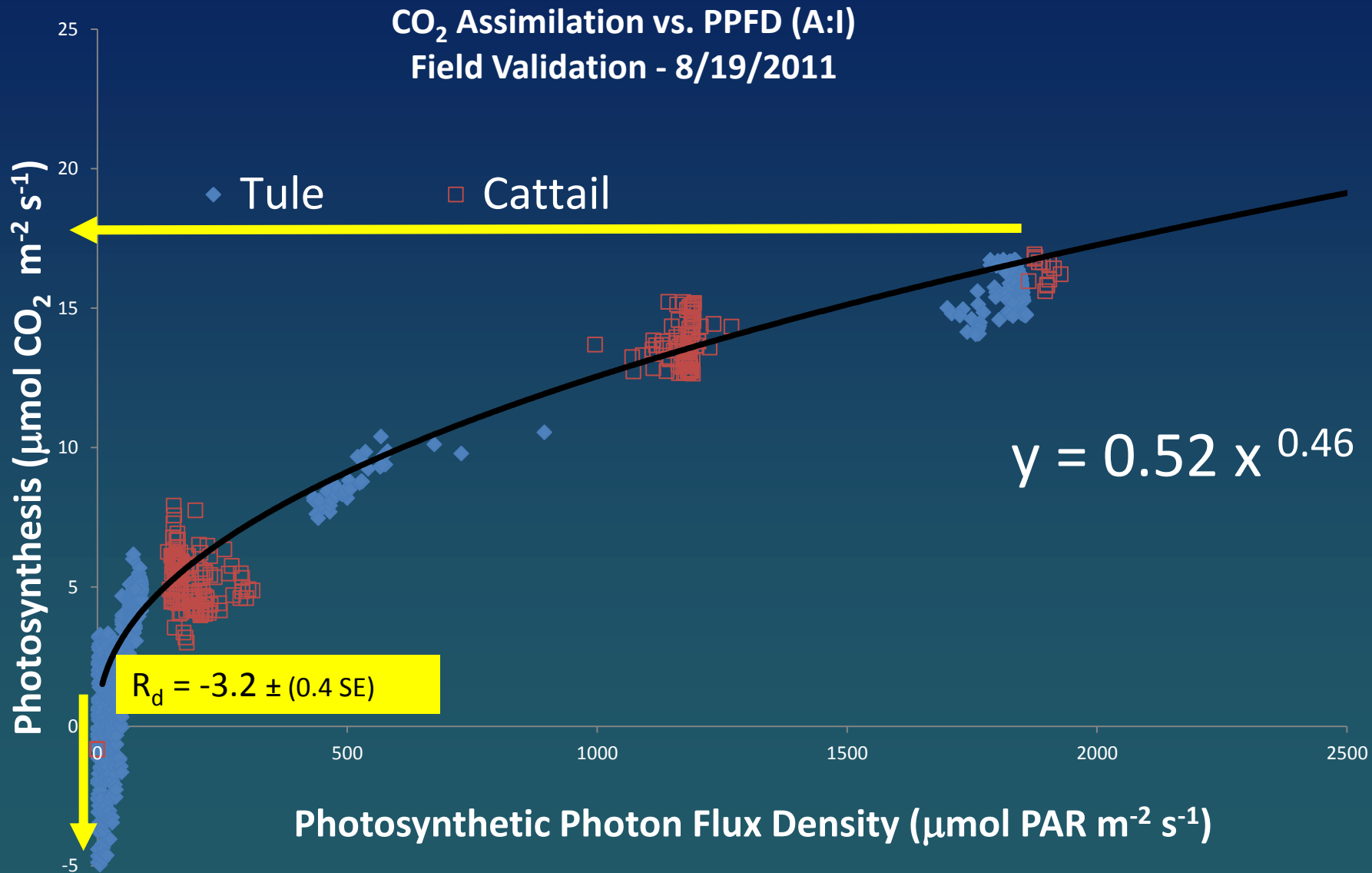
- LiCor XT6400 (light and CO₂ control)
- A:I and A:CI Curves (link to annual PAR)
- Field Validation (May-September)
- 3 stations (A, B, and C)
- 3 species (Tule, Cattails (TYAN, TYLA))

RESPIRATION

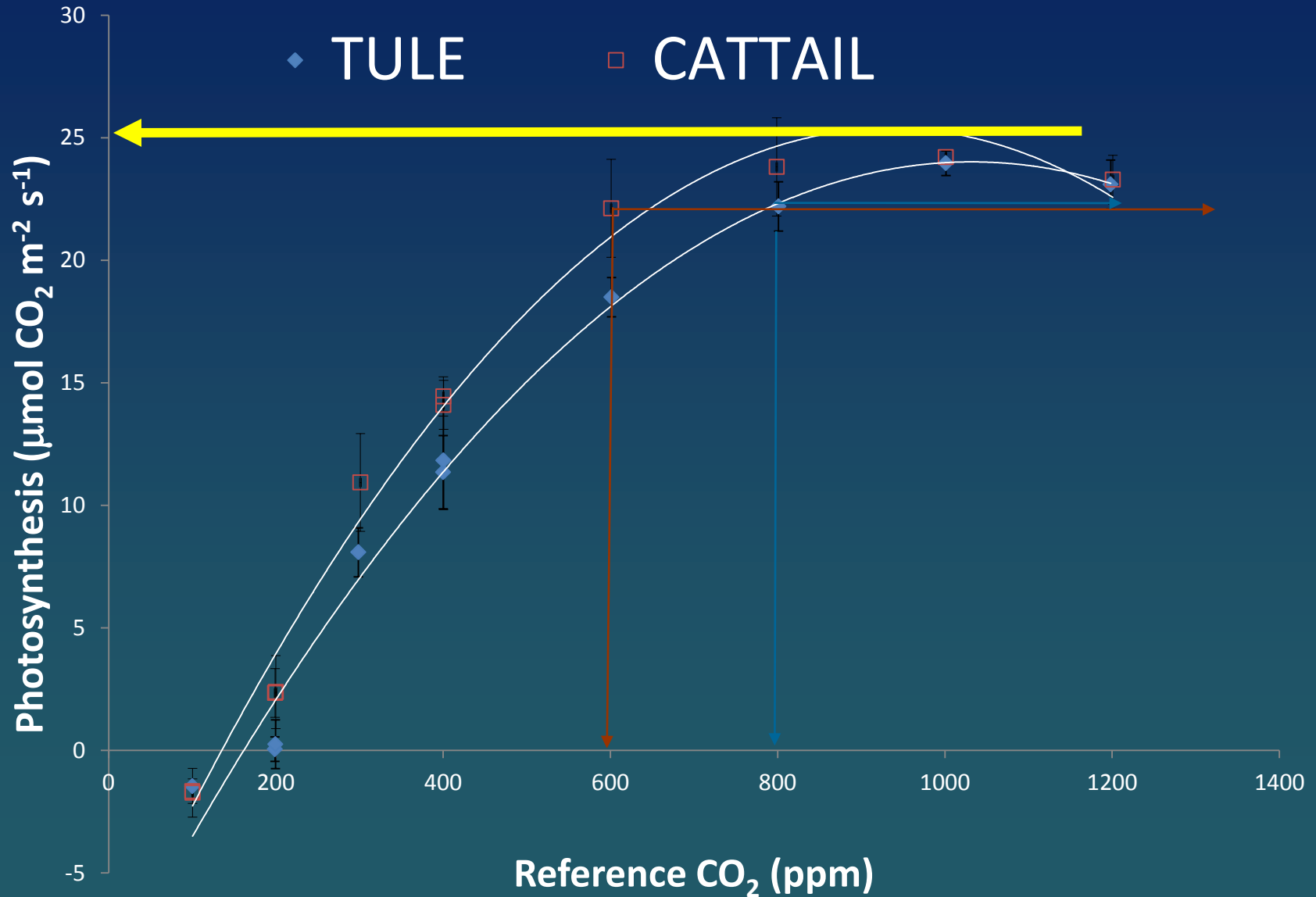
- Root Fermentation (Ethanol)
- Porewater acetate (and other Volatile Fatty Acids)
- Standing Biomass (NPP) and LAI
- Allometric $f(\text{density, height, diameter})$



PAR strongly controlled leaf photosynthesis



Response to elevated CO₂ : evidence for use of recycled CO₂

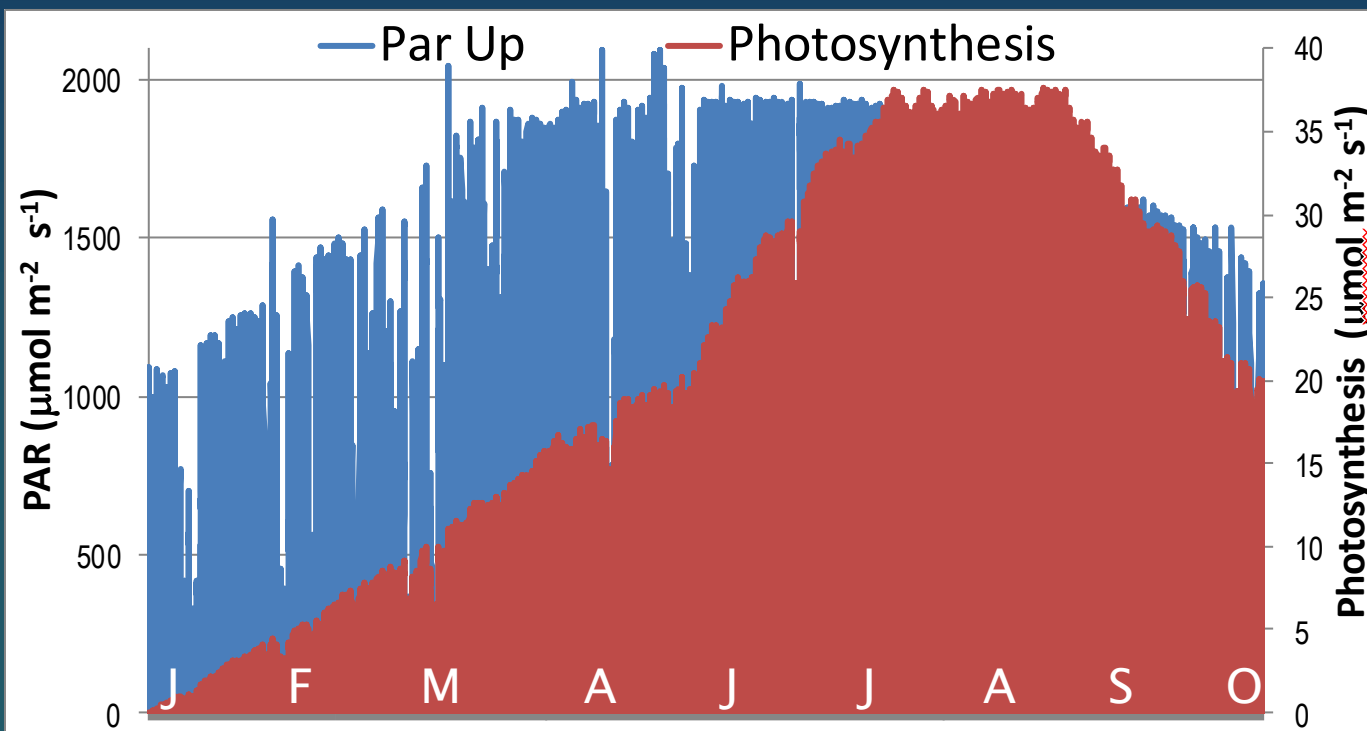


Carbon Use Efficiency: Annual GPP/NPP at Plot Scale (kg)

AGPP = Annual gross photosynthesis
 = (Photosynthesis x Curve-fit LAI)
 = $0.5231 (\text{PAR})^{0.46} \times \text{LAI}$

ANPP = Annual net biomass production
 = allometric biomass (Miller and Fujii 2009)
 = biomass x turnover rate

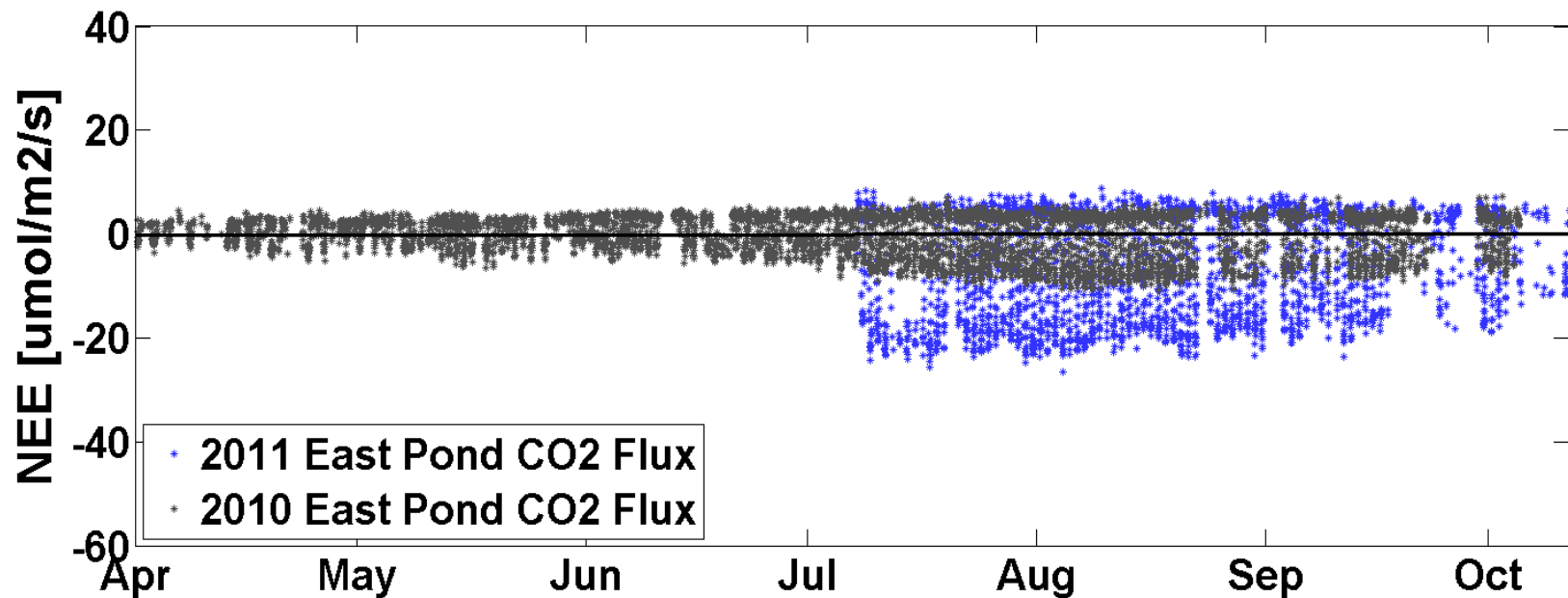
Station	Peak LAI	ANPP	GPP	CUE
A	6	7	9.9	0.70
B	2.4	4.4	6.8	0.63
C	2.2	3.2	4.9	0.68



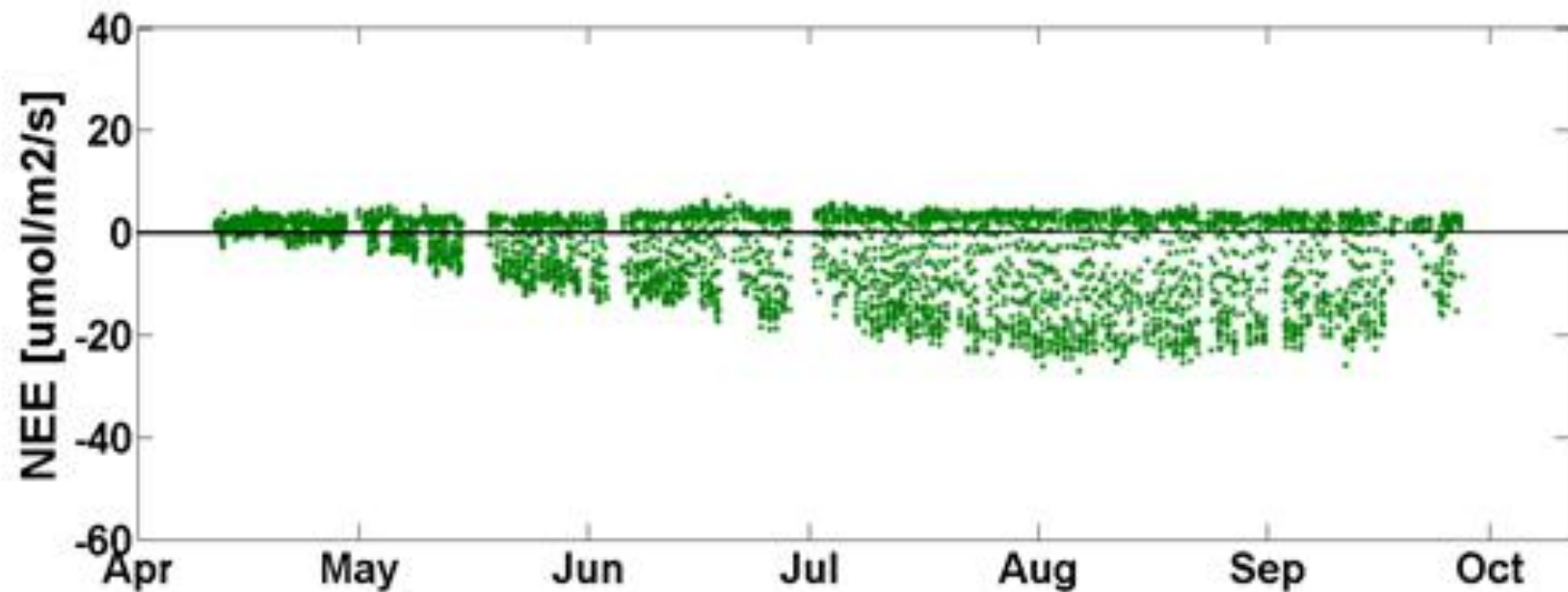
Kristin Byrd in Tule Thatch



East Pond CO2 Flux

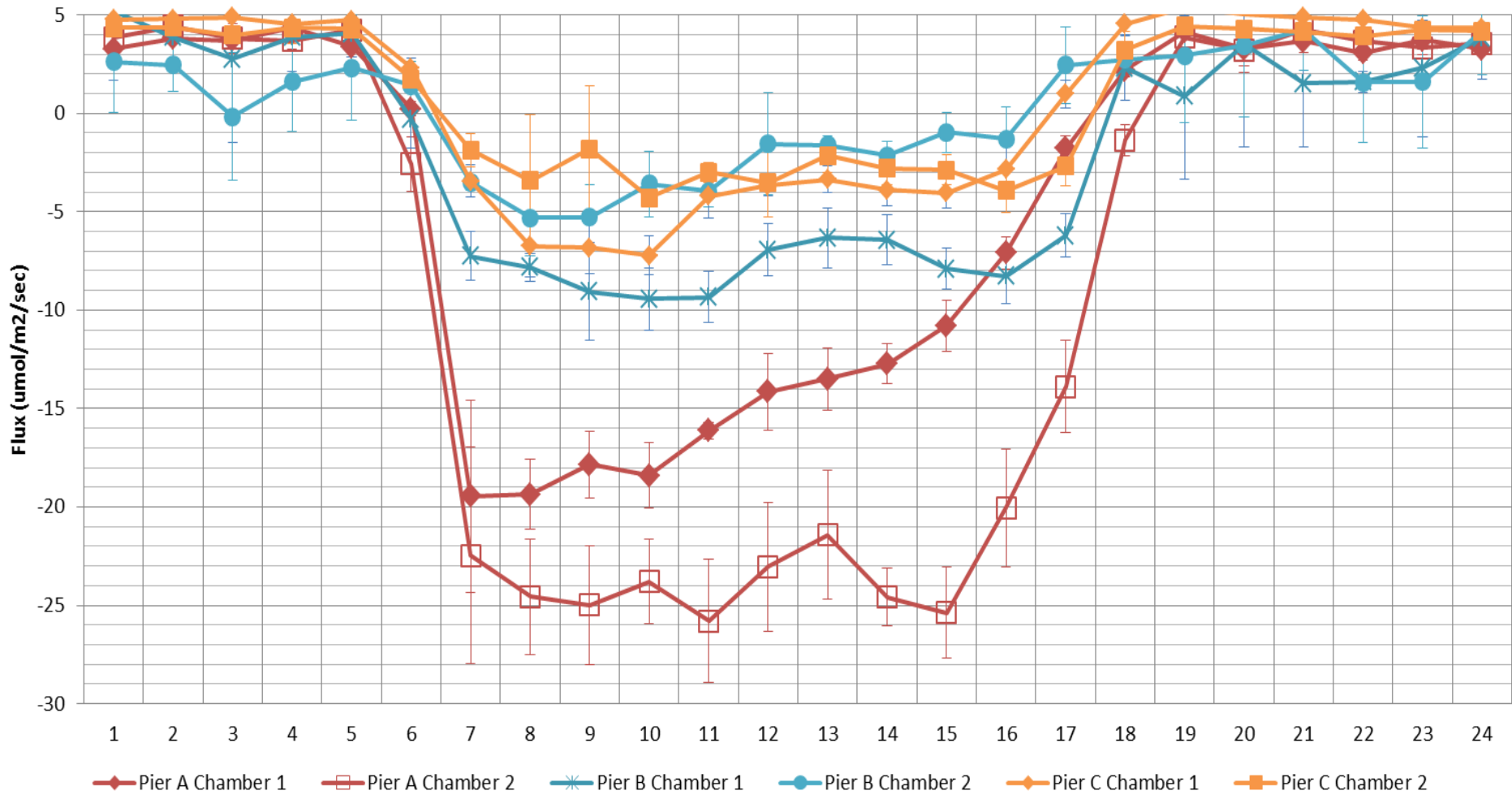


2011 West Pond CO2 Flux



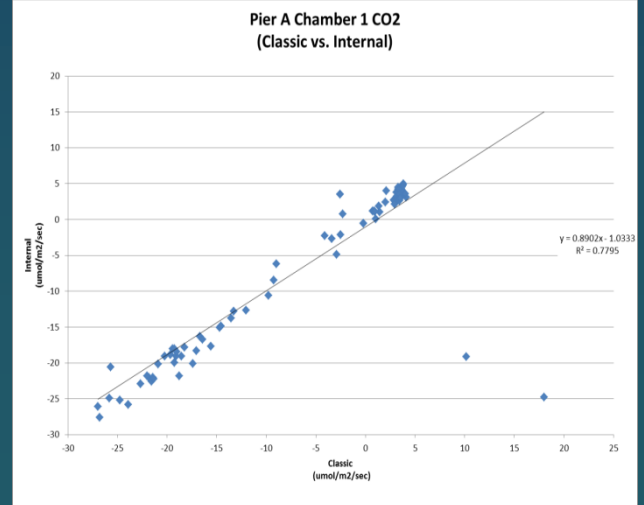
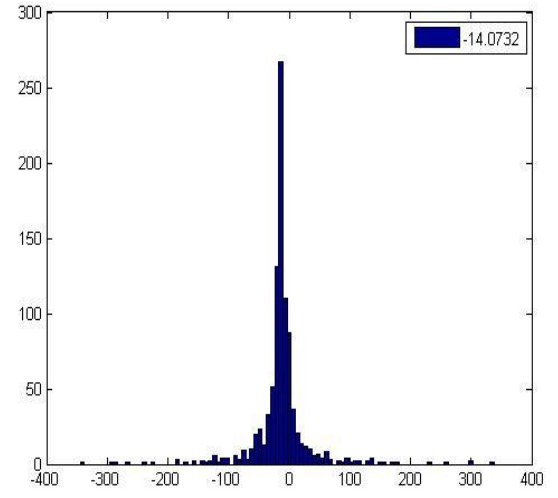
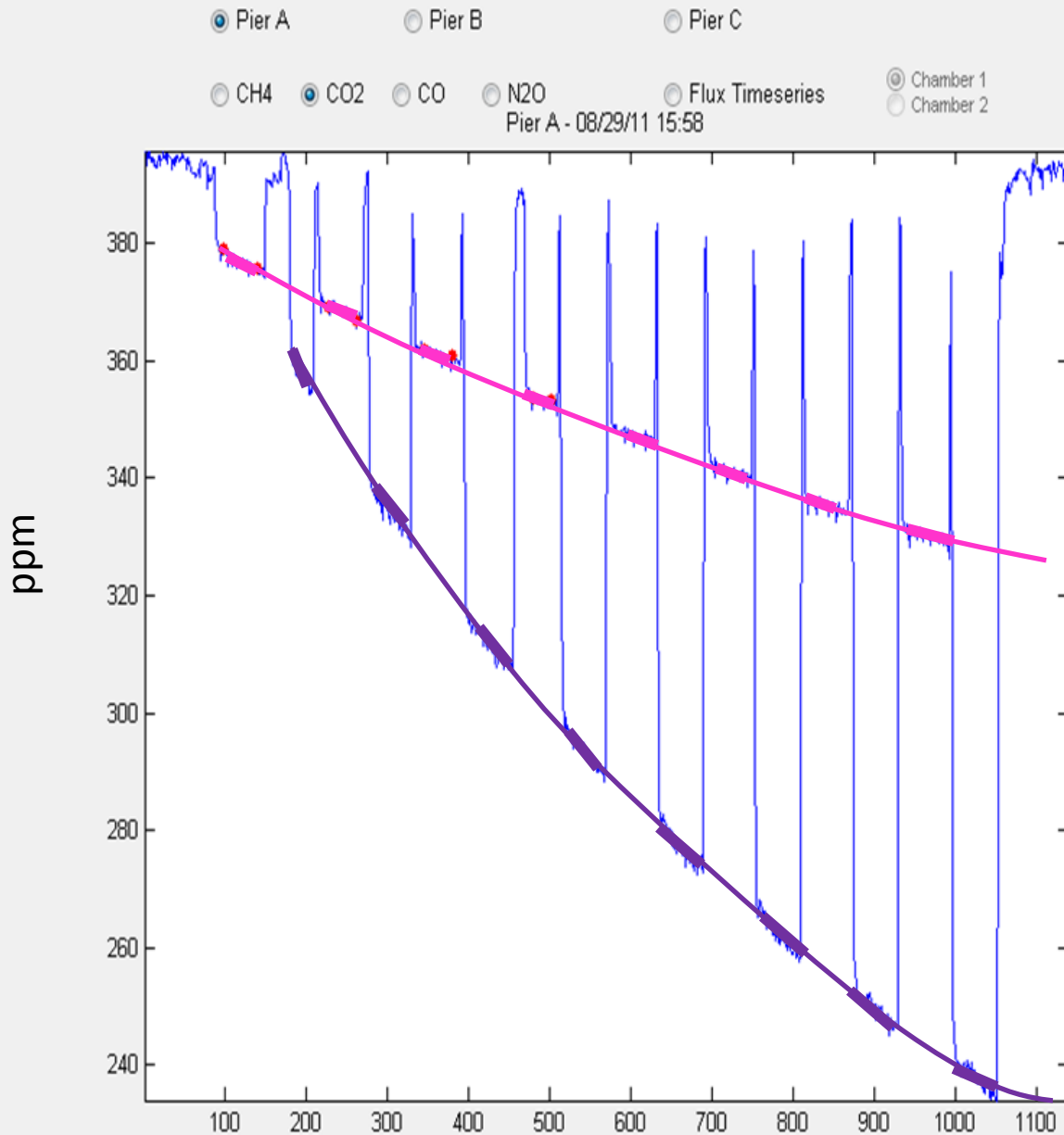
Static Chamber NEE

Mean Daily CO2 Flux - Piers A, B and C



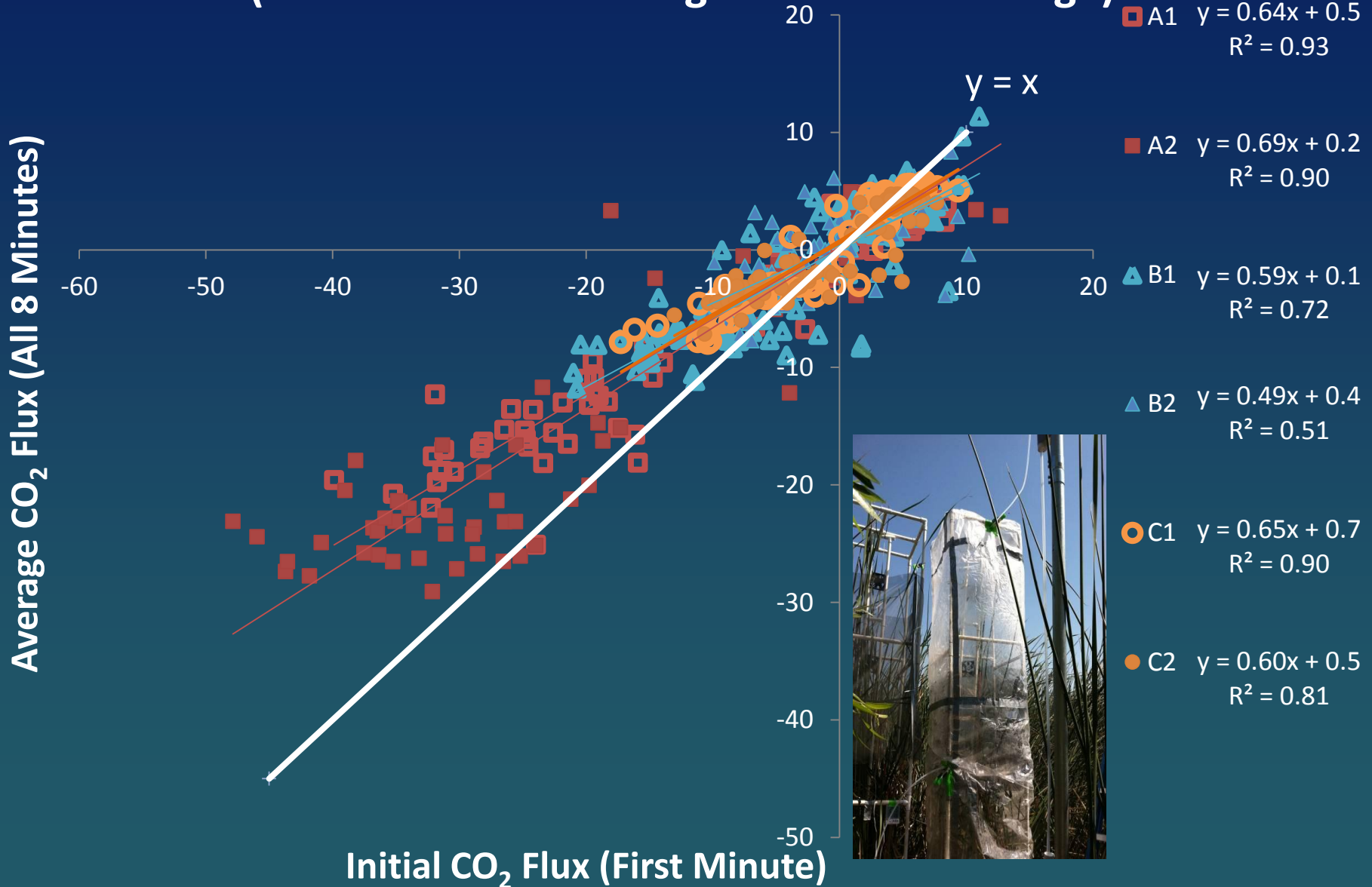
Static Chamber NEE Calculations

- Classic (8 median points)
- Classic (all points)
- Internal (handselected)
- Internal (median slope)



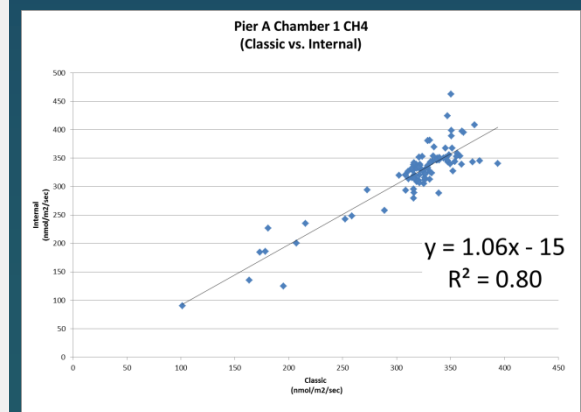
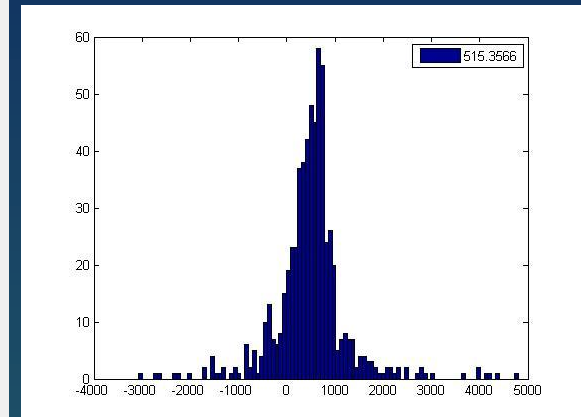
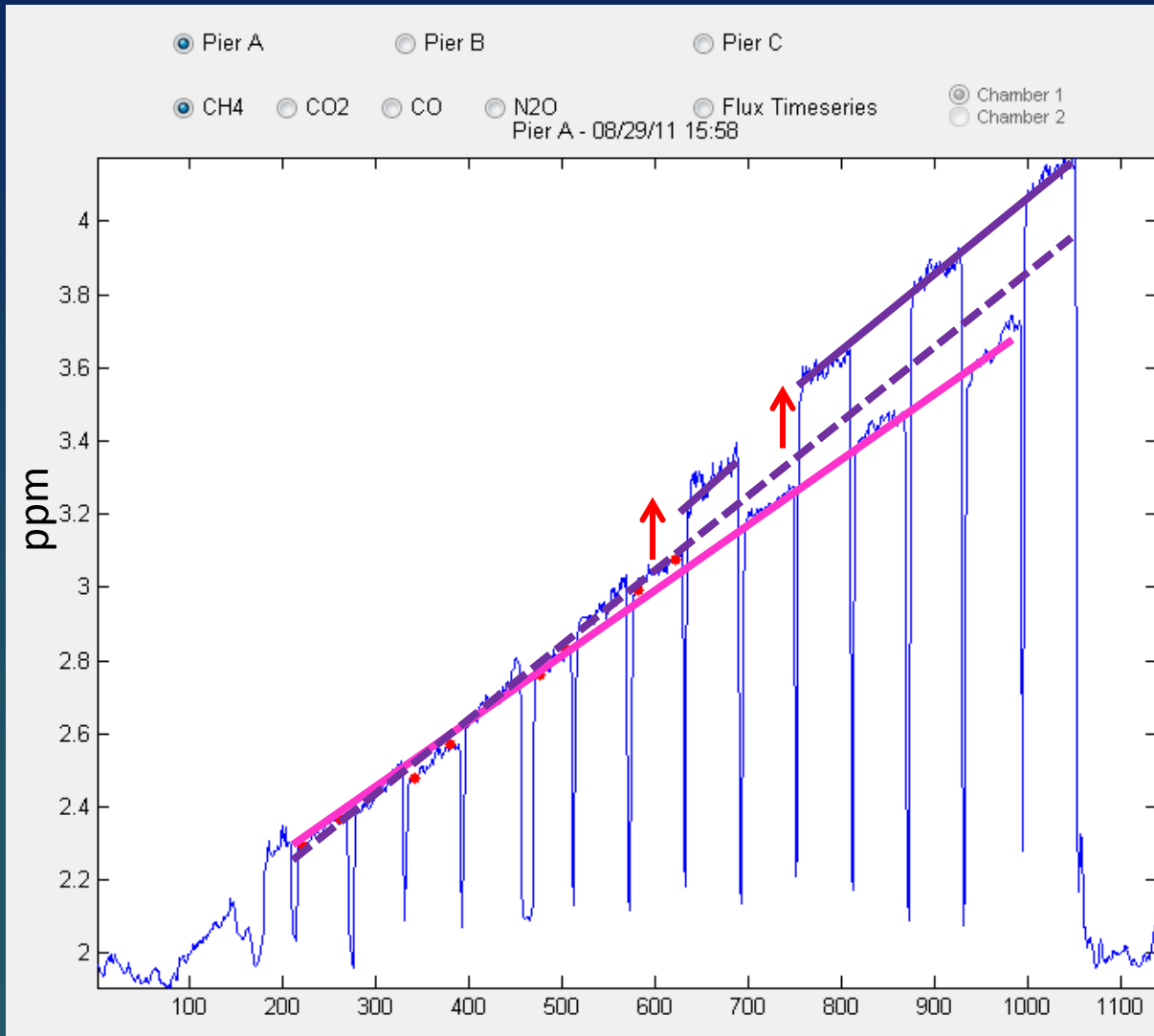
Chamber Effect

(first minute is 1.5-2x greater than average)

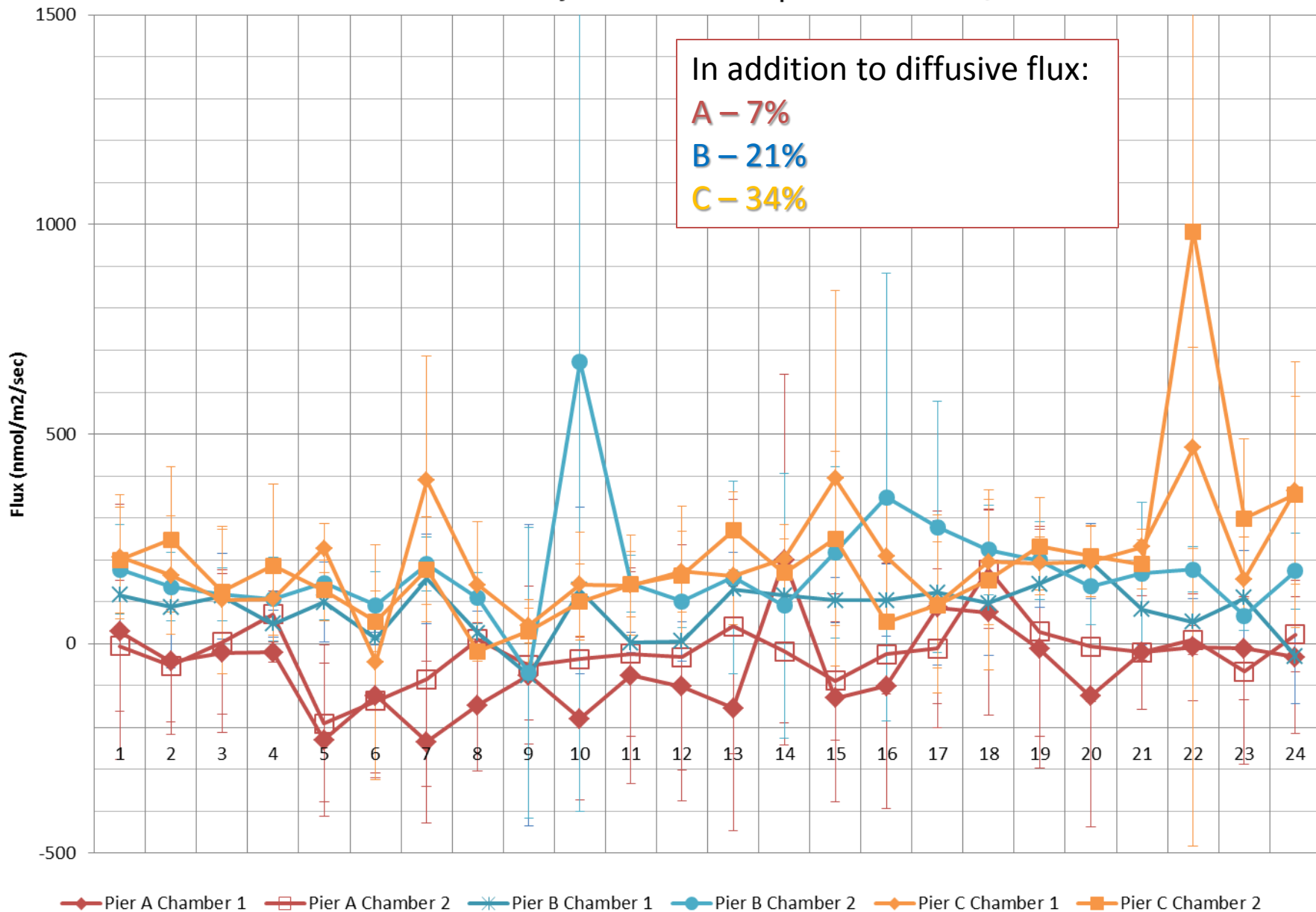


Static Chamber CH₄ Ebullition

- Classic (8 points)
- Classic (all points)
- Internal (handselected)
- Internal (median slope)

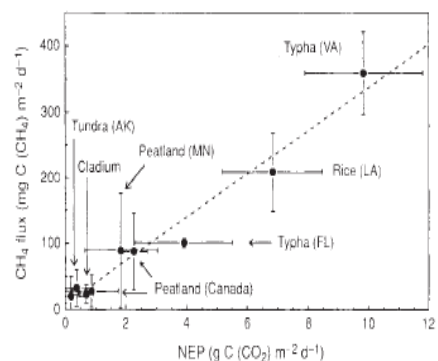


Mean Hourly Ebullition CH₄ Flux - Piers A, B and C



Spatial Differences (August 2011)

	C	B	A
mg CO ₂ m ⁻² d ⁻¹	-0.2	-15	-49
Spp/Stem Density	TYAN (79)	TYAN (41)	SCAC (226)
Abvgr. Biomass (kg)	2.0	1.8	4.4
Water Depth (cm)	17cm	17cm	20cm
Water Temp (Hi/Lo)	18/16	18/16	20/18
mg CH ₄ m ⁻² d ⁻¹	0.7	0.5	0.4
GWP mg CO ₂ eq m ⁻² d ⁻¹	+17	-3	-40



NPP vs CH₄
Whiting and Chanton 1993

C

B

A



METHOD COMPARISON (August 2011)

$\text{mg CO}_2 \text{ m}^{-2} \text{ d}^{-1}$

	C	B	A
Chamber	-0.2	-15	-49
Photosynthesis	-35	-38	-95
ECFlux	-15	-15	-15
$\text{mg CH}_4 \text{ m}^{-2} \text{ d}^{-1}$	0.1	0.1	0.1
GWP	-12	-12	-12
$\text{mg CH}_4 \text{ m}^{-2} \text{ d}^{-1}$	0.7	0.5	0.4
GWP	+17	-3	-40

C

B

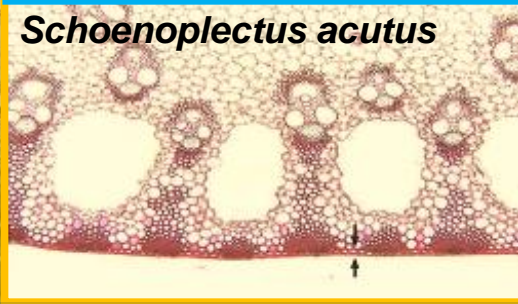
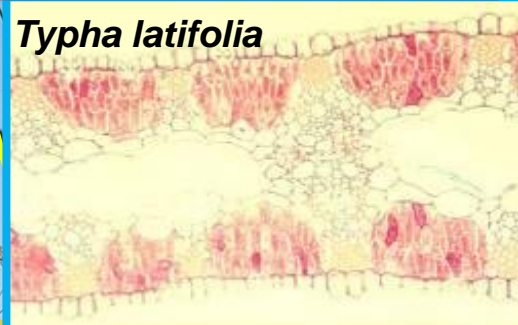
A



Leaf structure

Stomate size and density

Oxidized rhizosphere



1) THE PLANTS

- a) Moderate GPP w/high vertical LAI and low respiration
- b) Tule: Lignified, aerenchymous, high stomatal density

2) Delta Breeze – Limit to nighttime respiration?

Cycle of a Delta breeze

The breeze is part of a circulation of air that is driven by two tendencies of nature: Sunlight causes temperatures to rise more quickly over land than water, and air flows from areas of high pressure to areas of low pressure.

Cool water creates high pressure

A. Air near the sea surface cools and sinks. This creates the high-pressure marine layer.

B. The marine layer goes vertically and may reach a height of 2,000 feet.

C. The higher air cools the layer, the further inland it can spread when low pressure forms over the Central Valley. Coastal clouds may be swept inland in the marine layer.

Hot land creates low pressure

1. In the morning, warming land begins to heat the air above it.
2. The heated, lighter air rises then cools and falls in a cycle that grows over time. This is called a convection current.
3. By afternoon, the hot air that goes vertically eventually high pressure will force it to spread horizontally.
4. The rising and spreading air creates a low pressure system close to the land.
5. The hot air travels out to sea, becomes denser and falls.
6. Completing the cycle, the Delta breeze is called Valley's baby.

At night as the land cools, pressure increases, ending Delta breezes.

Minimum pressure around 5 p.m.
Maximum pressure around 5 a.m.

The Sacramento Bee
SATURDAY
September 11, 1999

Source: Cecil Cunningham and David Pavia, National Weather Service
See graphics? Scott Flinch and Sean McEwen

Cool Pacific's gift to Valley summers

By John D. Cox
Bee Staff Writer

The Mediterranean has the Mistrals, North Africa its Sirocos, Southern California its Santa Anas, and California's Central Valley has its Delta breezes.

Make that, its beloved Delta breeze. Toward the end of some of the hottest days of summer, about the time that the morning's idea of a barbecue is feeling like a terrible mistake, along comes the gust of cool ocean air like a gift from the gods.

Temperature plummet. Humidification evaporates.

Altitudes improve. Data crawl out from flower beds. With the sense of relief comes high hopes for a cooler night and a following day that will be at least less intense.

"It's like nature's air conditioner," observes Scott Cunningham, a National Weather Service meteorologist in Sacramento whose job it is to forecast when the Delta breeze will reach into the Sacramento and San Joaquin valleys — and when it won't.

"Will the breeze reach Sacramento tomorrow? Will it make it only to Fairfield? As far as Davis? Will it come

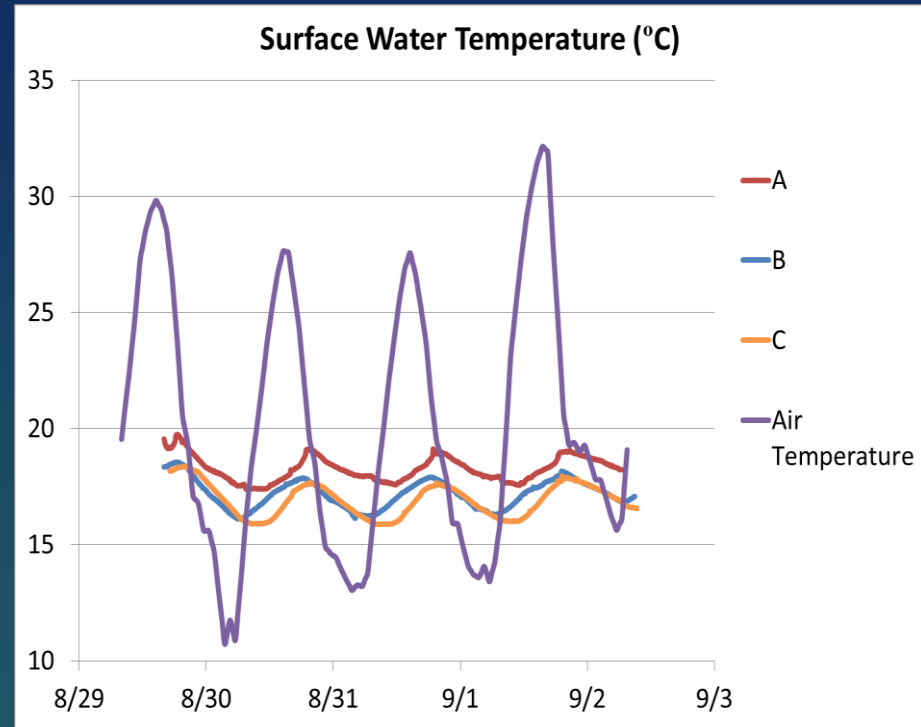
Please see BREEZE, page B4

Delta breezes



The Delta breeze cycle starts with rising and falling air currents. This happens in the Valley when mid-atmospheric air (about 20,000 feet) is light enough to allow lower air to rise.

Warm days (>30°C)

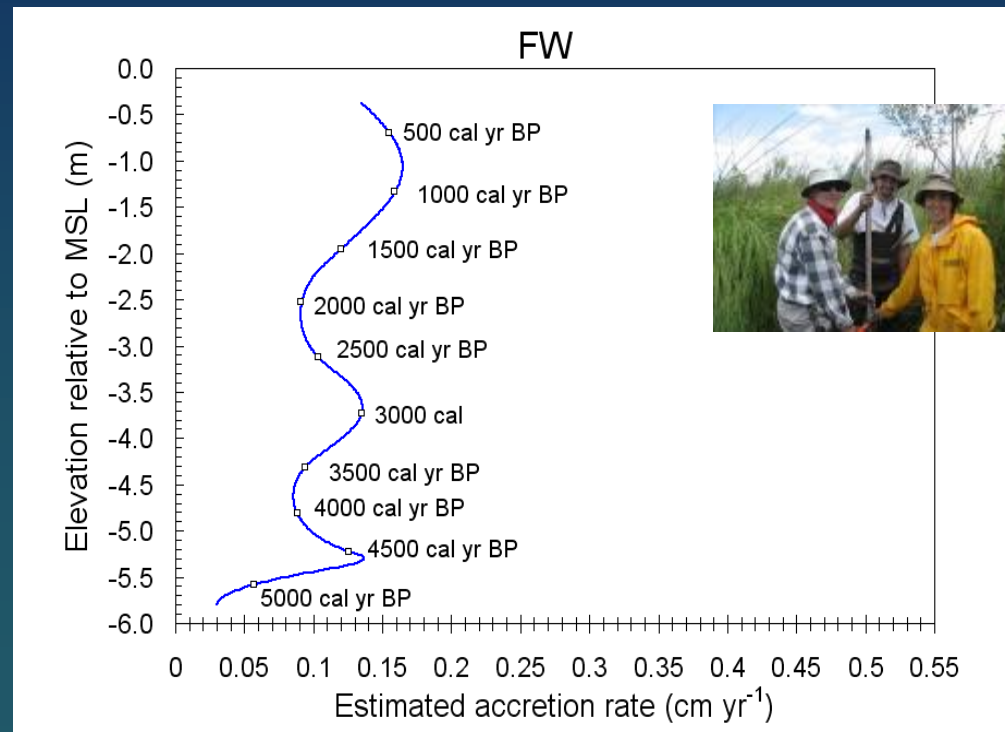


Cool nights (<15°C)

3) Constant, large accomodation space

“not your normal freshwater tidal wetland”

Freshwater tidal wetlands accrete peat at average rates of only $\sim 1.5 \text{ mm yr}^{-1}$ (RSLR)



Method Pro's and Con's

EC Flux -

Large footprint

Annual flux

Cannot separate CO₂-flux processes

Cannot separate CH₄-flux processes

Chamber-

Good spatial variability

CH₄-flux separated into diffusion or ebullition

Annual flux must be modeled

Leaf Photosynthesis -

Direct calculation of GPP and stress response

Process-based separation of CO₂-flux

Annual flux requires modeling w/ PAR and LAI

A sunset scene with a bright sun low on the horizon, casting a warm orange glow. Silhouettes of tall grasses and trees are visible against the bright sky.

Thank you.

lwindham@usgs.gov